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Measurements by controlled meteorological balloons in coastal areas of Antarctica

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Abstract: An experiment applying controlled meteorological (CMET) balloons near the coast of Dronning Maud Land, Antarctica, in January 2013 is described. Two balloons were airborne for 60 and 106 hours with trajectory lengths of 885.8 km and 2367.4 km, respectively. The balloons carried out multiple controlled soundings on the atmospheric pressure, temperature and humidity up to 3.3 km. Wind speed and direction were derived from the balloon drift. Observations were compared with radiosonde sounding profiles from the Halley Research Station, and applied in evaluating simulations carried out with the weather research and forecasting (WRF) mesoscale atmospheric model. The most interesting feature detected by the CMET balloons was a mesoscale anticyclone over the Weddell Sea and the coastal zone, which was reproduced by the WRF model with reduced intensity. The modelled wind speed was up to 10 m s^{-1} slower and the relative humidity was 20–40% higher than the observed values. However, over the study period the WRF results generally agreed with the observations. The results suggest that CMET balloons could be an interesting supplement to Antarctic atmospheric observations, particularly in the free troposphere.

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Key words: Dronning Maud Land, observations, troposphere, WRF

Introduction

Observations on the Antarctic atmosphere are sparse and mostly limited to the lowermost few metres of the atmosphere. Near-surface air temperature (T), relative humidity (RH), wind speed (ws) and direction (wd), and atmospheric pressure (P) are measured at manned and automated stations (Turner & Pendlebury 2004). Other meteorological activities at manned stations include visual cloud observations (at practically all stations) and measurement of solar shortwave radiation (Pirazzini 2004), thermal longwave radiation (Walden *et al.* 1998) and turbulent fluxes of momentum and heat (Handorf *et al.* 1999), as well as atmospheric composition (Barbaro *et al.* 2015). However, at many stations measurements are only made during the summer. Most of the manned stations are located in coastal regions and there are large gaps between stations, especially in West Antarctica. Automatic weather stations provide important additional near-surface meteorological observations in Antarctica (Van den Broeke *et al.* 2004, Lazzara *et al.* 2012). Higher in the atmosphere, radiosonde soundings measure the vertical profiles of T, RH and wind up to 20–30 km

(Durre *et al.* 2006). However, there are only about a dozen regularly operating sounding stations in Antarctica, most of them at coastal regions, and soundings are only made once or twice a day at most (Nygård *et al.* 2013). Over the Southern Ocean, the availability of *in situ* atmospheric observations is even more limited. In summer, a relatively small number of near-surface observations are made from ships, but only some of these utilize radiosondes. Marine meteorological observations have only been made during a few campaigns during the winter months (e.g. Andreas *et al.* 2002, Jonassen *et al.* 2015).

The available observations yield a basic understanding of weather and climate over Antarctica (King & Turner 1997), and provide the necessary initial conditions for numerical weather prediction models (Bromwich *et al.* 2005). Increasing the spatial and temporal resolution of atmospheric observations is needed for several reasons. First, Antarctic weather forecasting would benefit substantially from additional observations of atmospheric P and vertical profiles of T, RH and wind (Powers *et al.* 2012), with P and T observations being the most essential as initial conditions for operational



Fig. 1a. Controlled meteorological (CMET) balloon launched from Troll station, Antarctica, January 2012. **b.** CMET balloon just after launch. Photographs: P.B. Voss.

forecasting models. Second, more observations are needed to better evaluate the performance of numerical weather prediction and climate models in the Antarctic

(Tastula & Vihma 2011). Identification of the model errors and the conditions when they occur is a prerequisite for model improvements. Third, more *in situ* atmospheric observations are needed to provide ground-truth data for validation and further improvement of satellite remote sensing algorithms. Fourth, better observational coverage is needed for several specific research areas addressing, among others, atmospheric dynamics, circumpolar heat, moisture and mass budgets, atmosphere–ocean and atmosphere–snow/ice interactions, as well as climate variability and change. Climatological studies in Antarctica are currently subject to considerable uncertainties due to the sparsity of observations (Bromwich *et al.* 2014).

Opportunities to increase the spatial and temporal coverage of *in situ* observations from the Antarctic atmosphere are relatively limited. Increasing the number of radiosonde stations and the frequency of soundings would improve weather forecasts and provide valuable data for process studies and climatological research, but the high costs and the infrastructure required mean that a large increase in sounding activity appears improbable. The recent development of unmanned aerial vehicles (UAV) shows promise (Cassano 2014, Knuth & Cassano 2014, Jonassen *et al.* 2015) but the cost-effective operation of UAVs over large areas on a regular basis is still a major challenge. Hence, there is a need to identify new possibilities for cost-effective observation of the Antarctic atmosphere. As one possibility, we present data from a study in which two controlled meteorological (CMET) balloons (Voss *et al.* 2013, Stenmark *et al.* 2014) were launched near the coast of Dronning Maud Land in January 2013 (flying for 2.5 days and 4.5 days respectively). We show how the data can be used to evaluate simulations utilizing the weather research and forecasting (WRF) mesoscale atmospheric model (Skamarock *et al.* 2008).

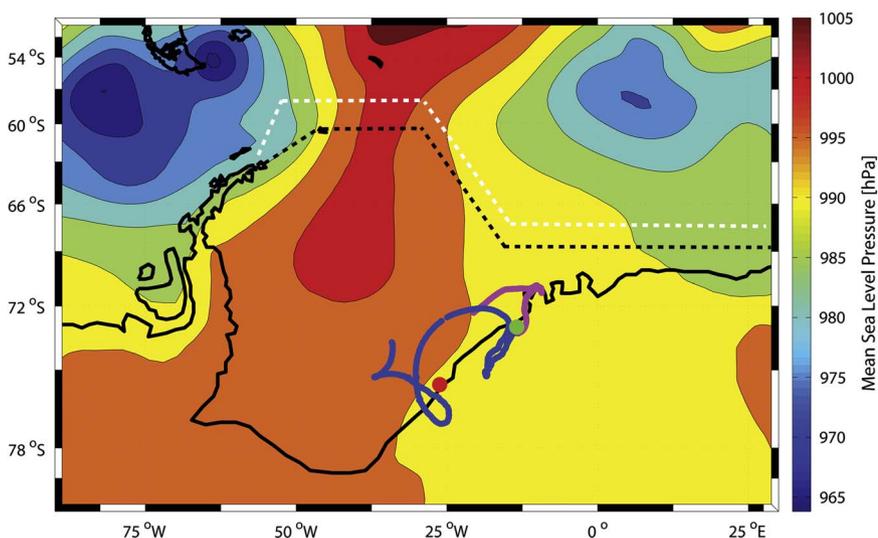


Fig. 2. Mean sea level pressure from the ERA Interim 80 km reanalysis (apps. ecmwf.int) for 19 January 2013 06h00 UTC. The Alpha balloon trajectory is shown in magenta, while the Bravo balloon trajectory is shown in blue. The green dot is the location of the launch site at Aboa (73.05°S, 13.41°W), while the red dot is Halley (75.61°S, 26.18°W). The white dashed line is the approximate envelope of the January 2013 ice edge, while the black dashed line is the 1979–2012 median January ice edge provided by the National Snow and Ice Data Center (nsidc.org).

Table I. Key characteristics of the controlled meteorological (CMET) balloon flights.

	Launch date and time	End date and time	Maximum altitude (m a.s.l.)	Latitude range (°S)	Longitude range (°W)
Alpha	16 Jan 2013 12h31 UTC	19 Jan 2013 01h40 UTC	3048	70.92 to 73.34	9.25 to 20.54
Bravo	18 Jan 2013 16h51 UTC	23 Jan 2013 02h01 UTC	3314	72.12 to 77.14	13.41 to 37.04

Observations

Controlled meteorological balloons

Controlled meteorological balloons have been used previously in both the Arctic (Roberts *et al.* 2015) and Antarctic (Stenmark *et al.* 2014) (Fig. 1a). The payload weight is *c.* 250 g and the standard CMET balloon consists of an outer zero-pressure balloon (ambient pressure) and an inner super-pressure balloon (0–60 kPa above ambient) (Fig. 1b; Voss *et al.* 2013). To make the CMET balloon descend, a small diaphragm pump transfers helium from the zero-pressure balloon into the pressurized balloon, thus reducing the volume of the entire system and increasing its density. To ascend, the process is reversed using a small valve to release helium and inflate the zero-pressure balloon. The CMET balloon will then rise or fall until its overall density (including the payload mass) matches that of the surrounding atmosphere. The payload includes a GPS receiver and satellite modem for data transfer and in-flight balloon altitude control, and sensors for T, P and RH. The balloon is powered by a solar panel and a small lithium polymer battery (see Voss *et al.* 2013 for details).

On 16 and 18 January 2013, two CMET balloons, hereafter called Alpha and Bravo, were launched from the Finnish Aboa station in Dronning Maud Land (73°03'S, 13°25'W; Fig. 2). The balloons were airborne for *c.* 60 and 106 hours, respectively (Table I). A third balloon failed after a short flight because of overfilling with helium. On 21 January 2013 at 22h00 UTC, Bravo started to descend from *c.* 3200 m a.s.l., whilst at a horizontal distance of only 0.7 km from the British Halley Research Station, located on the Brunt Ice Shelf off the Caird Coast, Dronning Maud Land, at 75.61°S 26.18°W, 30 m a.s.l. (Fig. 2). On 22 January 00h00 UTC, Bravo was at 378 m a.s.l. and 1 km from Halley. Consequently, this descent is well suited for comparisons with Halley radiosonde observations.

Radiosonde data

High-resolution radiosonde data from Halley for the study period were obtained from the British Antarctic Survey (BAS). The radiosondes are launched from Halley daily at 11h00 UTC and with an ascent speed of *c.* 3 m s⁻¹ they burst at *c.* 12 hPa 30 000 m a.s.l. after *c.* 150 minutes. The standard radiosonde measurements include P, T and RH, with ws and wd derived from the flight. The vertical resolution of the radiosonde data is 10–11 m.

Model experiment

The WRF model (Skamarock *et al.* 2008) version 3.5.1 was run in a three-domain set-up of 27, 9 and 3 km resolution (not shown). The model was initialized on 16 January 2013 at 00h00 UTC and was run for 192 hours until 24 January 00h00 UTC. The rapid radiative transfer model (RRTMG) (Iacono *et al.* 2008) was used for longwave and shortwave radiation schemes. The atmospheric surface layer was parametrized applying the Monin-Obukhov scheme, while the surface parametrization was based on a National Oceanic and Atmospheric Administration (NOAA) land-surface scheme operating with four layers including snow cover (Chen & Dudhia 2001). For the planetary boundary layer, the Yonsei University (YSU) turbulence scheme was utilized (Hong *et al.* 2004). Lastly, the Kain-Fritsch scheme was used for the cumulus parametrization (Kain 2004), which allows for deep and shallow convection. The YSU turbulence scheme has previously proven to work well in Antarctica (Tastula *et al.* 2012, Stenmark *et al.* 2014). The initial and boundary conditions were taken from the NOAA global forecast system (GFS) with 0.5 degree resolution.

Results

The observation period was characterized by a large anticyclone building up east of the Antarctic Peninsula from 18 January 2013 (Bravo launch) (Fig. 2).

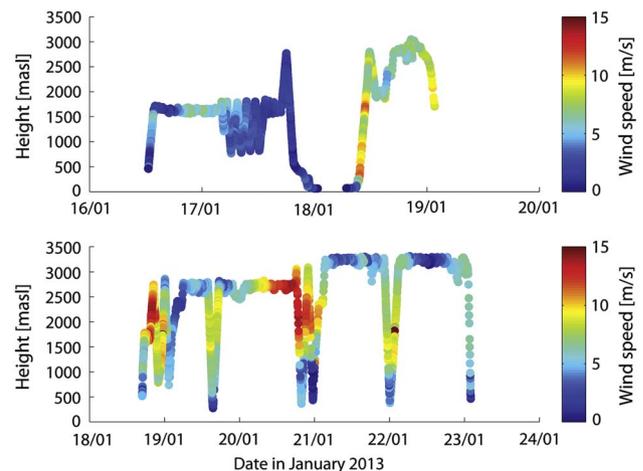


Fig. 3. Height profiles of Alpha (top) and Bravo (bottom), colour code is for wind speed (m s⁻¹). Note that the time axis is shifted between the panels.

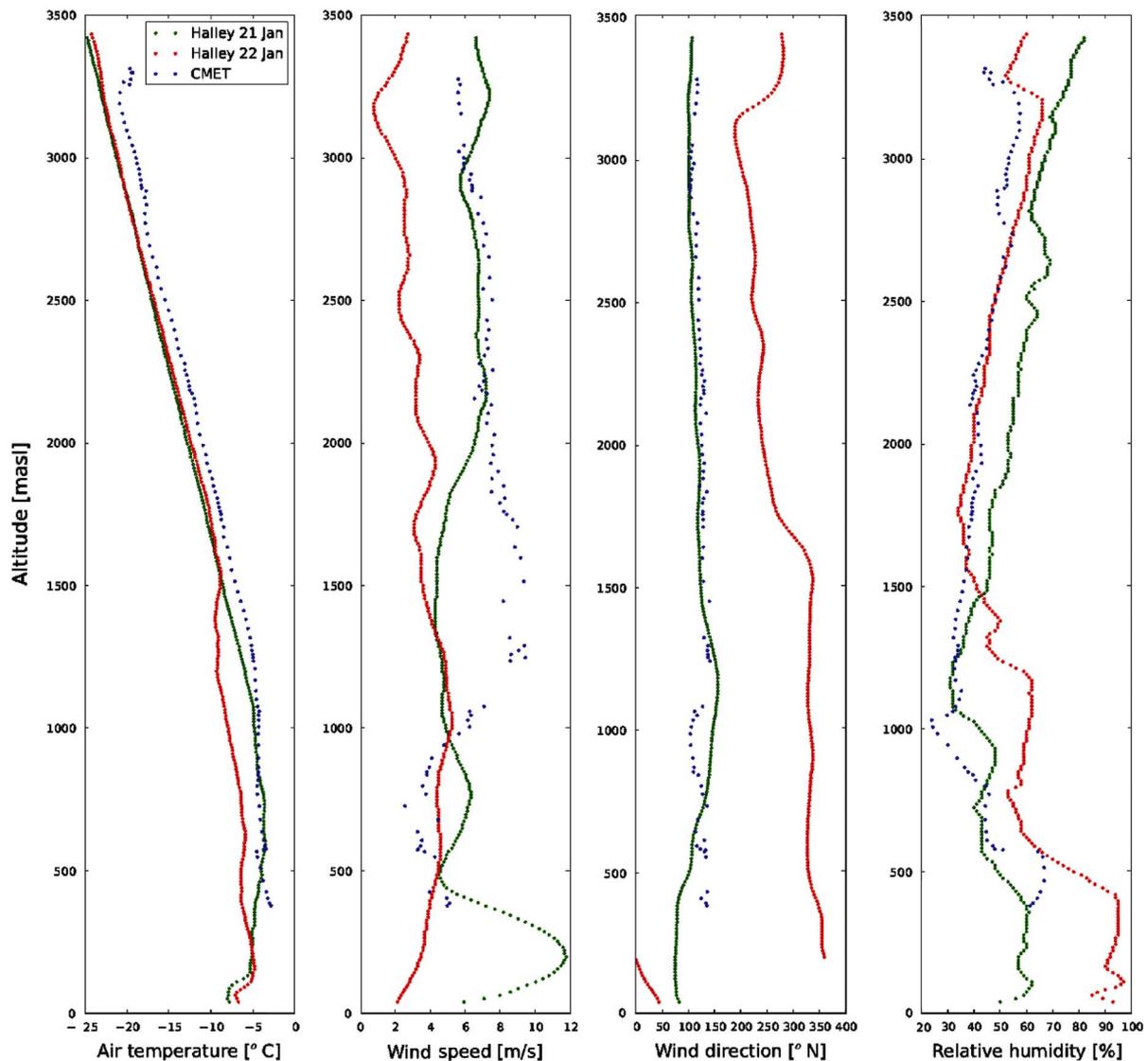


Fig. 4. Halley radiosonde profiles on 21 and 22 January 2013 at 12h00 UTC compared to nearby Bravo profiles on 22 January at 00h00 UTC.

Further east there was cyclonic activity centred *c.* 400 km off the Princess Astrid Coast on 19–21 January. Figure 2 shows the trajectories of the two CMET balloons. Alpha drifted north-east, until it turned westward over the ice shelf and continued over the Weddell Sea. Winds carried Bravo first to the south-west, but then it returned to Aboa along almost the same path. Then Bravo's trajectory showed a counter-clockwise semicircle over the Weddell Sea. The radius of the circle was *c.* 250 km, indicating a mesoscale anticyclone. Close to the end of the semicircle, Bravo passed Halley and continued over the continent on the Caird Coast. Then Bravo turned north-west and winds carried it over the Weddell Sea, where it dropped. The horizontal distances covered were 835 km for Alpha and 1989 km for Bravo. Figure 3 shows the altitude profiles of

the two balloons. Alpha carried out multiple shallow soundings between 750 and 1750 m a.s.l. on day 2 of the flight, while Bravo carried out several deeper soundings between 300 and 3200 m a.s.l., particularly on days 3 and 4.

Figure 4 compares the profiles of *T*, *ws*, *wd* and *RH* from the CMET balloons and the Halley radiosondes on 21 and 22 January 2013. The radiosonde profiles were taken *c.* 12 hours before and after the CMET balloon descent profile. The *T* measurements compare reasonably well, although with a slight warm offset.

Since the CMET balloons are moving in three dimensions with time, it is not appropriate to only compare them with vertical model profiles. Instead a four-dimensional approach was used, where each observation from the CMET balloons was compared

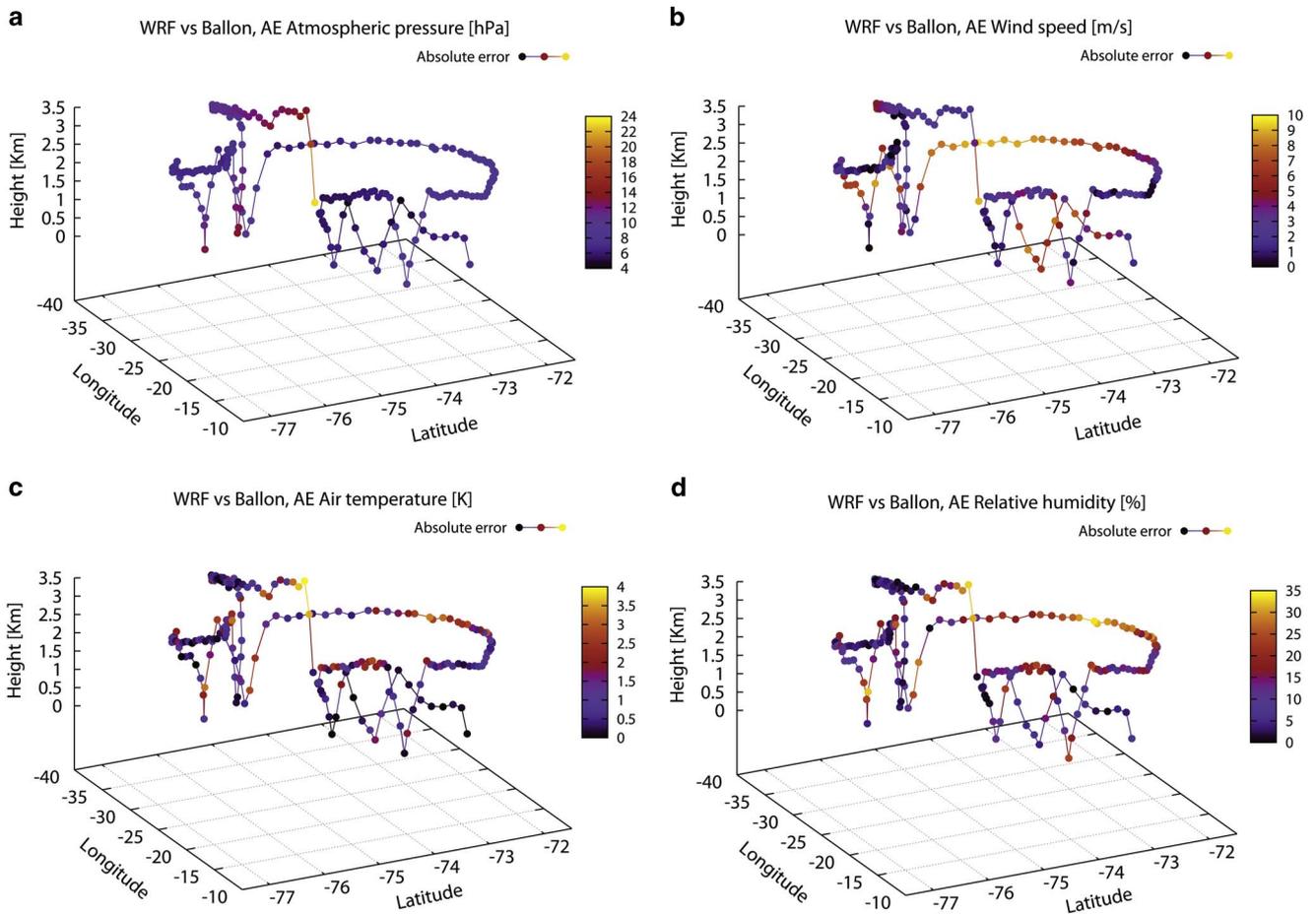


Fig. 5. 4-D comparison of controlled meteorological (CMET) balloon observations and weather research and forecasting (WRF) model output for the nearest grid point for each model time step (30 min). Output is shown for **a.** air pressure, **b.** wind speed, **c.** air temperature and **d.** relative humidity.

with the nearest model grid point in time and space. Figure 5 shows the mean absolute errors (MAE) for P, ws, T and RH from the Bravo flight (also summarized in Table II for Alpha and Bravo). These results indicate that the statistical model performance is independent of the choice of turbulence scheme. Comparing Figs 2 and 5 we see that the largest MAE in ws, T and RH occur when

Bravo is cruising westward at *c.* 2500 m a.s.l. over the Weddell Sea on 20 January.

Discussion and conclusions

The most interesting feature detected by the CMET balloons was the mesoscale anticyclone over the Weddell Sea and the coastal zone. Mesoscale anticyclones are common in the ocean (e.g. Sabu *et al.* 2015) but much less in the atmosphere. The cases reported include the following. Blamey & Reason (2009) detected an offshore mesoscale anticyclone at 800–700 hPa levels along the east coast of South Africa, and concluded via model experiments that coastal orography played an important role in its generation. Nishikawa *et al.* (2014) observed a mesoscale anticyclone over the Kuril Islands between the North Pacific and the Sea of Okhotsk. They concluded that it was generated by an extremely cold spot in the sea surface temperature. Mesoscale anticyclones have also occurred in Antarctic coastal regions but have received very little attention, as mesoscale cyclones have been more of a focus (Heinemann & Klein 2003).

Table II. Comparison of weather research and forecasting (WRF) model output and controlled meteorological (CMET) balloon observations for relative humidity (RH), air temperature (T), dew point temperature (Td), wind speed (ws) and air pressure (P). The mean absolute errors are given for WRF outputs applying the Yonsei University (YSU) and the Mellor-Yamada (MYJ) turbulence schemes.

	RH (%)	T (K)	Td (K)	ws (m s ⁻¹)	P (hPa)
Alpha					
YSU	15.9	1.5	6.2	2.1	9.1
MYJ	15.0	1.5	5.9	2.4	9.2
Bravo					
YSU	11.6	1.0	3.1	3.1	8.2
MYJ	12.1	1.1	3.3	3.1	8.3

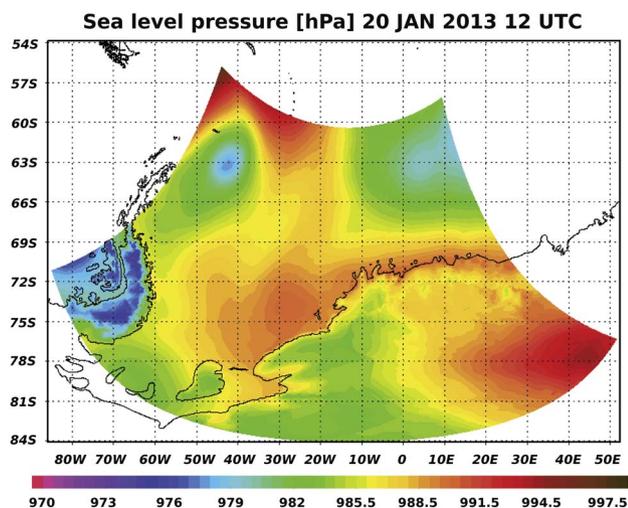


Fig. 6. Weather research and forecasting (WRF) sea level pressure in the inner (3 km) domain on 20 January 2013 at 12h00 UTC.

In our case, the mesoscale anticyclone was reproduced by the WRF experiment (Fig. 6), but it was less intense than the actual observations. The modelled w_s was up to 10 m s^{-1} lower and the RH was 20–40% higher than the values observed by Bravo (Fig. 5). The low w_s in the model suggests that the pressure gradient related to the anticyclone was not strong enough in the model. However, over the study period, the WRF results generally agreed with the CMET balloon observations (Fig. 5) and radiosonde soundings at Halley (Fig. 4). The WRF output was also compared to the Halley weather station observations (not shown), and the w_s values were within $1\text{--}2 \text{ m s}^{-1}$. The WRF model generated mesoscale cyclones slightly north–north-west of Halley on 18–23 January (not shown), but these were not resolved by the ERA Interim analysis. The sudden change in Bravo's trajectory west of Halley and the consecutive counter-clockwise drift (Fig. 2) suggests that there was indeed a mesoscale cyclone in the region. Furthermore, other loops and sudden turns in the trajectories of Alpha and Bravo suggest the presence of mesoscale flow features, but the timings and locations do not match with the WRF results.

According to Heinemann & Klein (2003), challenges in modelling mesoscale processes in the Antarctic coastal zone include parameterization of subgrid-scale physical processes, representation of the coastal orography and the inaccuracy of initial conditions. More recently, Zhang *et al.* (2015) addressed the problems in modelling the wind field over Antarctic coastal orography. All of these modelling challenges were present in our study. In addition, during the study period the sea ice coverage in the Weddell Sea was more extensive than average, with the ice edge *c.* 200–300 km further north (Fig. 2), which may have influenced cyclogenesis in the region. The anomalous sea ice cover was not well represented in the

GFS analysis used for the boundary conditions in the WRF model, which may have contributed to the differences in the observed and simulated cyclones.

To overcome the modelling problems more observations are needed for i) better identification of the errors, ii) further development of the subgrid-scale physical parameterizations, and iii) data assimilation to operational models. The CMET balloons have a lot of potential to yield observations for these purposes. In addition, as CMET balloon data are Lagrangian (except when vertical profiles are measured by adjusting the buoyancy of the system), they are well suited for mass budget studies of the Antarctic atmosphere. The mass flux associated with katabatic winds is assumed to be significant even on a global scale (Dalu *et al.* 1993). The mass lost by the katabatic outflow must be replaced by subsidence over Antarctica (Van de Berg *et al.* 2007) which in turn is fed in the middle and upper troposphere by convergence of air into a cyclonic vortex above Antarctica (Starr & Vedder 1989, King & Turner 1997). A larger number of CMET balloons could yield valuable information on these processes, or at least contribute to the evaluation of atmospheric reanalyses and model products.

In the future, a suite of CMET balloons could be launched with a strategy of having some long-duration CMET flights to follow air masses at constant altitude, and for other CMETs to undertake repeated soundings-on-command to profile the atmosphere at specific times and locations of interest. Future development plans for the balloons include adding a drag string or proximity sensor so that soundings can reliably descend to within a few metres of the surface without damaging the sensor payload. The potential ability to conduct soundings in remote and otherwise inaccessible regions (such as crevasse areas) at low cost is rather unique to CMET balloons. Some UAVs may be capable of such measurements, but their operation requires significantly more personnel time, making long-term or repeated measurements less feasible. Repeated CMET soundings, into katabatic winds for example, could improve meteorological model performance in Antarctica yielding better forecasts and simulations of mass balance of the Antarctic ice sheet. Furthermore, the lifetime of the CMETs can almost certainly be increased (to 20 days or more) by using better balloon films or supplemental ballast. The dry climate in Antarctica is optimal for lightweight airborne sensors, and the problem of icing that is common to cold regions is avoided here. In addition to being able to carry out controlled soundings, the CMET balloons can collect long time-series at fixed altitudes (isentropic surfaces) in the free troposphere. This could be valuable for model validation and for tracking air masses, for example during atmospheric chemistry studies (Voss *et al.* 2010).

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Author contribution statement

Dr Lars R. Hole was involved in the field measurements and was primarily responsible for data analysis and writing the manuscript. Mr Alexis Bello was responsible for WRF simulations and analysis, comparisons with observations and figure preparation. Dr Tjarda Roberts and Dr Timo Vihma were involved in discussing the results and putting them into context. Dr Paul Voss was responsible for field measurements and analysis of observations.

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